Analysis of Stream-Temperature Variations in the Upper Delaware River Basin, New York

GEOLOGICAL SURVEY WATER-SUPPLY PAPER 1999-K

Prepared in cooperation with the Delaware River Basin Commission



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By OWEN O. WILLIAMS

CONTRIBUTIONS TO THE HYDROLOGY OF THE UNITED STATES

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UNITED STATES DEPARTMENT OF THE INTERIOR ROGERS C. B. MORTON, Secretary

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CONTRIBUTIONS TO THE HYDROLOGY OF THE UNITED STATES

ANALYSIS OF STREAM-TEMPERATURE VARIATIONS IN THE UPPER DELAWARE RIVER BASIN, NEW YORK

By OWEN O. WILLIAMS

ABSTRACT

The effect of climatologic conditions and reservoir releases on downstram conditions was determined by means of statistical and graphical analyses of stream-temperature variations measured in the upper Delaware River besin, May-September 1964-67. Climatologic conditions normally increase water temperatures from February through July and decrease them from August through January. Summer releases from New York City's Cannonsville Reservoir vere observed to decrease water temperatures by 13°C (Celsius) in 8.1 miles and by 1°C, 55.9 miles downstream from this reservoir. Releases from New York City's Pepacton Reservoir were observed to decrease water temperatures by 11°C in 31.0 miles and between 1°-3°C in 71.0 miles downstream from this reservoir. The influence of releases from these reservoirs is dependent upon five factors: Thermal stratification in the reservoir, depth at which water is withdrawn from the reservoir, rate of release, distance downstream from the reservoir, and climatologic conditions.

INTRODUCTION

In the spring of 1964, the U.S. Geological Survey and the Delaware River Basin Commission entered into a cooperative program to collect and report stream temperatures in the upper Delaware River basin. Collection of data was scheduled only during the summer months from 1964 through 1967. The objectives of this cooperative investigation were to: (1) Define the variations in water temperatures during the summer months, (2) determine the effect of reservoir releases on downstream water temperatures, and (3) provide data from which the thermal suitability of the study area as a spawning and nursery habitat for anadromous fish can be determined.

Stream temperatures always have been important to studies of water quality because many physical, chemical, and biological properties of water are expressed as functions of temperature. Their importance is more apparent now because of the increased awareness of the effect of thermal pollution upon ecology. This term "thermal

pollution" usually is applied to influxes of thermally overleaded water. However, an influx of thermally depleted water can also be a form of thermal pollution and often has a marked effect on the stream's ecological system. Plant or animal life, which may have flourished in an area, can become scarce or even be eliminated if temperature requirements are not met. At the same time, other plant or animal life, which may have been scarce, may begin to flourish.

Influxes of thermally depleted water can result from reservoir releases made through a low-level outlet during summer months. Most large storage reservoirs exhibit the classical type of thermal stratification. Mackenthun, Ingram, and Porges (1964, p. 9-10) describe thermal stratification as follows:

For a few weeks in the spring, water temperatures may be homogeneous from the top of a water body to the bottom. Vertical water density is also homogeneous and it becomes possible for the wind to mix the water in a lale*** * The advance of summer quickly checks circulation by warming the surface waters; as they warm they become lighter, resting over colder water of greater density. Thus, a permanent thermal stratification is formed. In natural deep bodies of water, three layers eventually form. The upper layer, or epilimnion, represents the warm, more or less freely circulating region of approximately uniform temperature, and may vary in thickness from 10 feet or less in shallow lakes to 40 feet or more in deeper ones. The middle layer, or thermocline, is the region of rapid change, usually defined by a change in temperature of 1.8° F (Fahrenheit) for each 3.28 feet variance in depth. The lower layer, or hypolimnion, is the cold region of approximately uniform depth. It is cut off from circulation with upper water.

As autumn comes, the standing body of water cools: the epilimnior increases in thickness until the lake becomes homothermous, and again a period of complete circulation begins. This occurs from late September to December, depending upon the area and depth of the lake and its geographic and climatic location. It lasts until changes in density reestablish stratification, or until the lake is frozen over. This commonly occurs from November to January, varying with lake and season and geographic location. Circulation then ceases until spring.

If a storage reservoir is stratified in this manner and releases are made through a low-level outlet from the hypolimnion, the water being released will be colder than the normal stream temperature in the area. If the release is both significantly colder and significantly larger than the natural flow in the stream, the water temperature downstream from the reservoir may be lowered so that the temperature requirements of certain plant or animal life in the stream are not met.

GENERAL FEATURES OF THE STUDY AREA

The reach under study (fig. 1) extends 17.5 miles from Cannonsville Reservoir on the West Branch Delaware River to its confluence with the East Branch at Hancock; from Pepacton Reservoir on the East Branch, 32.6 miles to the confluence with the West Branch; and 54.5

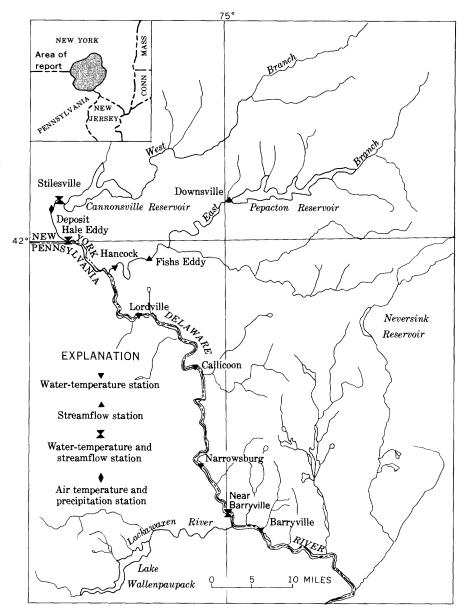


FIGURE 1.—Location of temperature-monitoring sites and streamflow-gaging stations.

miles along the main stem Delaware River from the confluence of the East and West Branches at Hancock to Barryville. The drainage area of the Delaware River at Barryville is 2,692 square miles. The drainage areas controlled by the Cannonsville and Pepacton Reservoirs are

454 and 371 square miles respectively. Pertinent information on mileages and drainage areas is presented in table 1.

Both Cannonsville and Pepacton Reservoirs are in the headwaters of the Delaware River, where the slopes of stream channels are steep. Each dam is more than 100 feet high, and water is released from a lowlevel outlet. The water in storage in both reservoirs spreads far beyond the former river channel into numerous fingers and embayments to provide a large surface area. Passage of water through the reservoirs may be discontinuous, and significant parts may remain in storage for nearly a year. Both reservoirs store substantial volumes of water during the winter and spring months and release substantial volumes during the summer and fall months, so as to maintain a minimal flow. Prior to completion of the Cannonsville Reservoir, a minimal flow of 1,525 cfs (cubic feet per second) in the Delaware River at Montague, N.J. (27.5 miles downstream from Barryville), was required. Upon completion of the Cannonsville Reservoir, the minimal flow to be maintained was increased to 1,750 cfs, subject to seasonal adjustment upward to 2,650 cfs, under a formula prescribed in a U.S. Supreme Court decree, amended in 1954.

Water-temperature data collected in the Pepacton Reservoir from 1959 to 1963, by the New York City Board of Water Supply, show that the reservoir becomes thermally stratified during the summer months. Because the Cannonsville Reservoir is a large storage reservoir in the same geographical area, there is little doubt it too becomes thermally stratified.

Table 1.—Downstream distances from New York City reservoir sites and drainage areas

	Distance, i	n miles—	Dusinose
Location	From Cannonsville Reservoir	From Pepacton Reservoir	Drainage area (sq mi)
Pepacton Reservoir		0	371
Cannonsville Reservoir	. 0		454
Stilesville			456
Hale Eddy	8. 1		593
Downsville		. 5	371
Fishs Eddy		21. 6	783
Hancock		31. 0	838
Confluence of two branches	17. 5	32. 6	1, 504
Lordville		41. 7	1, 587
Callicoon	44. 3	59.4	1,706
Narrowsburg	55. 9	71. 0	1, 913
Near Barryville		81. 6	2, 023
At Barryville		87. 1	2, 692

COLLECTION OF DATA

Three types of installations have been used by various agencies to collect temperature data in the upper Delaware River basin.

At Stilesville (fig. 1) a thermograph attachment to the water-stage recorder inside a shelter has been used since 1962 to record stream temperature and water stage continuously on the same strip chart. The sensing element, a pressure-sensitive type, was in an intake pipe on the bottom of the river about 15–25 feet from shore during low water. These records are collected by the Geological Survey in its continuing cooperative program with the city of New York. Selected daily maximum and minimum temperatures from this record are listed in table 4.

The New York City Board of Water Supply installed similar thermograph recorders at Hale Eddy and near Barryville (fig. 1) during the summer of 1963. The sensing element, a thermocouple, is in a fashion similar to that at Stilesville. These recorders were installed inside Geological Survey shelters as units separate from the water-stage recorders. Variations in water temperature are recorded on a weekly circular chart. Data were obtained seasonally during August and September 1963 and during May through September 1964–67. These records also are summarized in table 4.

Waterproof thermograph recorders were installed as part of this project in the spring of 1964 at Hancock on the East Branch Delaware River and at Lordville, Callicoon, Narrowsburg, and Barryville on the main stem of the Delaware River (fig. 1). The waterproof thermograph recorder is contained in a stainless-steel case with an O-ring seal. These recorders were chained to two 100-pound blocks and placed on the streambed in flowing water from 3 to 6 feet deep. They were in operation for approximately 5 months, May through September 1964–67. The instruments were serviced once every 20–25 days. All stream-temperature data are verified by the manufacturers to be accurate within 1° (Celsius). Daily-maximum and daily-minimum stream temperatures at these stations are summarized in table 4.

In addition, streamflow records were obtained at continuous-recording gaging stations at Stilesville, Hale Eddy, Downsville, Fishs Eddy, and near Barryville (fig. 1). Also, maximum and minimum air-temperature records at two sites in the upper Delaware River besin, Deposit and Port Jervis (19.1 miles down stream from Barryville), were obtained from the U.S. Weather Bureau (1963–67).

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ENVIRONMENTAL ANALYSES

Stream-temperature data collected in the upper Delaware River basin were analyzed to define: (1) Time and space variations and (2) the influence of climatologic conditions and reservoir releases upon stream temperatures. Statistical techniques used include frequency analyses, double-mass curve analyses, and simple graphical comparisons. Monthly statistics for each water-temperature monitoring site in the upper Delaware River basin are summarized in table 5.

VARIATIONS IN WATER TEMPERATURE

The relation between monthly mean temperature of the Delaware River at Narrowsburg, near Barryville, and at Barryville, and the monthly mean air temperatures at Deposit and Port Jervis (U.S. Weather Bureau, 1963-67) is depicted in figure 2 for the 1964-67 water years. Monthly mean stream temperatures for October through April during the period were estimated from records collected at Narrowsburg during 1948-51 (U.S. Geological Survey, 1954a, b, 1955). The plot shows that the monthly mean temperature of the Delaware Piver from Narrowsburg to Barryville is within 4°C of the monthly mean air temperature at Port Jervis, May through September 1964-67. Also, about 95 percent of the time, monthly mean stream temperatures at the three sites plot within 1°C of each other. The net difference in monthly mean stream temperature is ±1°C throughout the entire 16 mile reach. Indications are that stream temperatures relate more to climatologic conditions than to hydrologic conditions in this reach and that monthly mean air temperatures at Port Jervis are reasonable indices to the monthly mean stream temperatures from Narrowsburg to Barryville.

The use of seasonal air temperature as a reasonable index to seasonal water temperature was discussed for Pennsylvania stream by Mangan (1946, p. 9-10) and for Oregon by Moore (1967, p. 5-6). Mangan

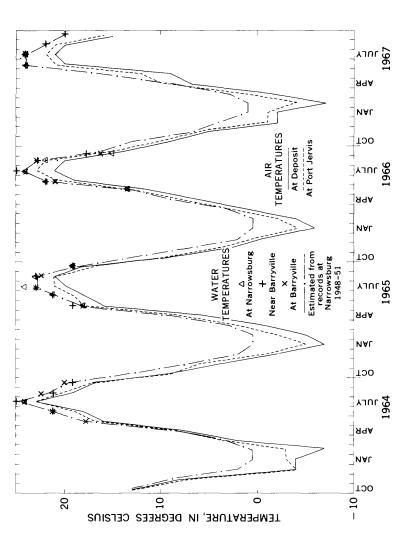


Figure 2.—Monthly mean water temperature in the upper Delaware River from Narrowsburg to Barryville and monthly mean air temperature at Deposit and Port Jervis, 1964-67 water years.

excepts streams that are affected by artifical impounding, and Moore also excepts streams that are either largely spring fed, north-south oriented, or shaded by wooded banks.

Data collected in the upper Delaware River between Narrowsburg and Barryville for this project indicate that long-term averages are not necessary for summer months, as the monthly mean water temperatures relate sufficiently well to monthly mean air temperature.

The Delaware River is basically north-south oriented. However, it is wide enough to avoid much shading by trees. Thus, water flowing in the reach from Narrowsburg to Barryville is exposed to direct sunlight during summer months. The river absorbs thermal energy at a slower rate than does the atmosphere, but also loses this energy at a much slower rate. This slower loss rate is apparently the reason that the monthly mean water temperatures tend to exceed the monthly mean air temperatures from March through November. During the winter months of December, January, and February, the lower the mal limit of 0°C also maintains the water temperature above the air temperature.

To define the seasonal variations of water temperature in the Delaware River from Narrowsburg to Barryville, the monthly mean water temperatures at three sites were averaged and a least-squares analysis made according to procedures described by Dixon and Massey (1957). These computations produced the following linear-regression equations:

Month	Equations
May	y=16.43+1.11 (x-16.00)
June	y=21.93+0.875 (x-20.00)
July	y=24.02+0.545 (x-22.25)
August	y=22.22+0.358 (x-21.25)
September	y=18.70+1.27 (x-16.25)

where y is the estimated monthly mean water temperature at any site between Narrowsburg and Barryville and x is the morthly mean air temperature at Port Jervis. Using these equations, the monthly mean water temperatures of the Delware River between Narrowsburg and Barryville can be estimated from the monthly mean air temperatures. A plot of the observed monthly mean water temperatures and monthly mean air temperatures and the estimated regression line for each of the above equations is shown in figure 3.

To test the validity of these equations in describing the seasonal variations in stream temperature of the Delaware River between Narrowsburg and Barryville, mean monthly air temperatures at Port Jervis were computed for May to September by averaging 5 years (1963–67) of air-temperature data (U.S. Weather Bureau, 1963–67). These average air temperatures were assumed to be the normal for each month and were used to estimate the corresponding mean monthly

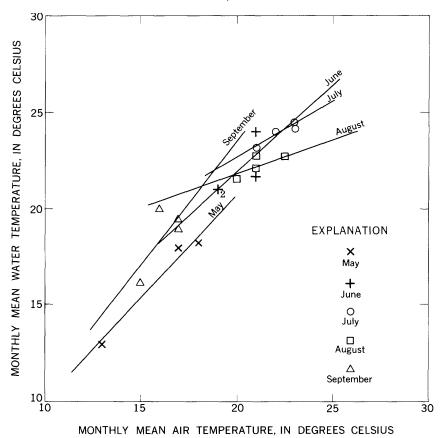


FIGURE 3.—Relation of monthly mean water and air temperatures from Narrowsburg to Barryville, May-September 1964-67. described by linear-regression analysis.

water temperatures for May through September from the equations on p. K8. The computed and observed mean monthly water temperatures were the same for each month: 16°C for May, 22°C for June, 24°C for July, 22°C for August, and 19°C for September.

Upstream from Narrowsburg, temperatures of the main stem of the Delaware River do not relate as well to air temperatures as temperatures at downstream sites; stream temperatures depart from air temperature even more in the East Branch and West Branch Delaware River. Figure 4 depicts the monthly mean water temperatures on the main stem above Narrowsburg and on the East and West Branches and depicts monthly mean air temperatures at Deposit for the periods May-September 1964-67. This plot shows that water temperatures generally increase in a downstream direction and that the monthly

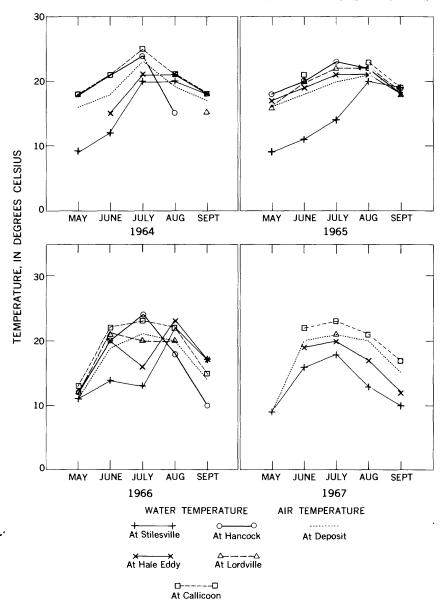


FIGURE 4.—Monthly mean water temperatures on the main stem Delaware River at Callicoon and Lordville, on East Branch Delaware River at Hancock, and on West Branch Delaware River at Hale Eddy and Stilesville, and monthly mean air temperatures at Deposit, May-September 1964-67.

mean water temperatures at Lordville and Callicoon were consistently higher than either those on the East Branch or those on the West Branch and, at times, higher than both.

The illustration shows that monthly mean water temperatures deviate increasingly from monthly mean air temperatures in an upstream direction. This deviation indicates that in the East and West Branches of the Delaware River a factor or factors other than climatologic conditions are affecting water temperature.

The influence of this factor is great enough to maintain lower stream temperatures than might be expected from local climatologic conditions. For example, the monthly mean water temperatures (fig. 4) at Stilesville, Hale Eddy, and Lordville are lower for July 1966 than for June 1966, whereas the monthly mean air temperature is higher.

INFLUENCE OF RESERVOIR DISCHARGES MONTHLY WATER TEMPERATURES

Monthly mean water temperatures on the West Branch Delaware River at Stilesville, prior to initiation of storage at and releases from Cannonsville Reservoir, and monthly mean air temperatures at Deposit and Port Jervis for May-September 1963 are depicted in figure 5. This illustration shows that the monthly mean water temperature at Stilesville generally related to the monthly mean air temperature at Deposit. However, figure 4 shows deviations in this relation for later data collected at this site.

A comparison of monthly mean air and water temperatures with monthly mean streamflow at Stilesville for the periods May-September 1963-67 (fig. 6) suggests that this lack of relation between air and water temperatures probably results from reservoir releases. The monthly mean streamflows at Stilesville prior to 1964 can be considered natural flow; streamflow subsequent to 1964 is a combination of releases and overflow from Cannonsville Reservoir. The mean streamflow at Stilesville for May 1963 was 566 cfs and the mean water temperature was 14°C (table 5). Monthly mean water temperature of the water released from the reservoir was colder during 1964-67 than the natural flow in May 1963. Monthly mean water temperatures for June and July during 1964-67 are also lower than those observed in June and July 1963 at Stilesville.

The effect of water released through the reservoir's low-level outlet is most noticeable in the plot for June and July 1966 in figure 6. Releases from Cannonsville Reservoir in July were much greater than those in June; water temperatures in July were lower than in June, although this occurred at a time when natural downstream temperatures would be expected to be increasing. This suggests that the large volume of water released in July influenced downstream temperatures to a greater extent than did the small volume released in June.

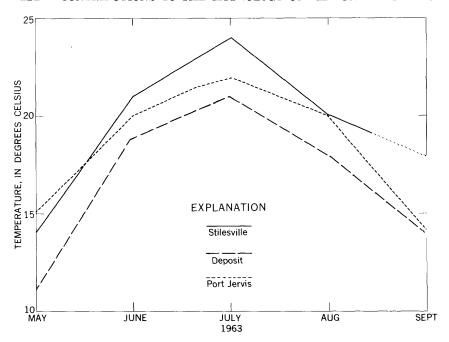
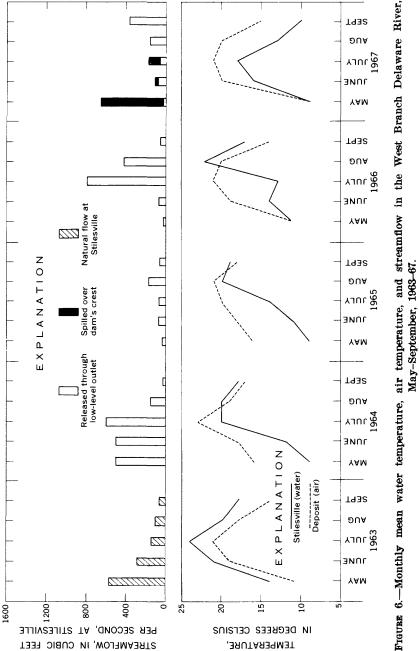
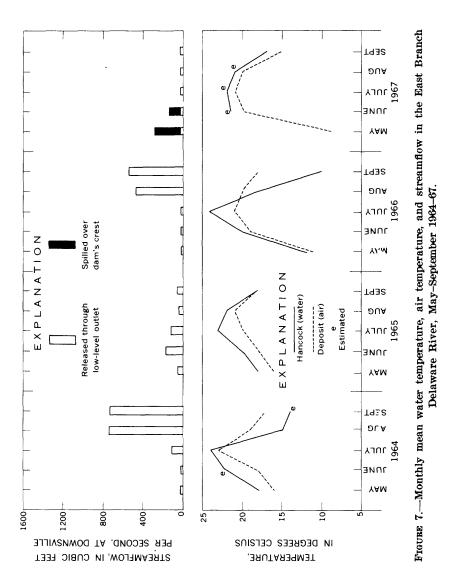


FIGURE 5.—Monthly mean temperature of the West Branch Delaware River at Stilesville and monthly mean air temperature at Deposit and Port Jervis, May—September 1963.

The relation between monthly mean air and water temperatures for August and September 1964–66 (fig. 6) approximates that for August and September 1963, but the August and September 1967 curves do not. A possible explanation is that releases through the low-level outlet from May to July in the 1964–66 period, at a time when the reservoir always stored less than 50 percent of its total capacity, effectively depleted the hypolimnion—the cold water in the bottom of the reservoir. Thus, by late summer in 1964–66, warmer water was being released. In 1967, however, the reservoir was filled to capacity and low-level releases were entirely from the hypolimnion through the summer and early fall.

A similar comparison of streamflow on the East Branch Delaware River at Downsville with water temperatures at Hancock and air temperatures at Deposit (fig. 7) indicated that the effect of releases from the Pepacton Reservoir on downstream water temperatures in the East Branch Delaware River is dependent upon the release rate. Streamflow data at Downsville, 0.5 mile below the reservoir, was used to indicate the reservoir-release rate. When releases from this reservoir were relatively small, as during May-September 1965, monthly mean water temperatures at Hancock (31 miles below the reservoir) were





as much as 3°C warmer than those of the air. When releases were relatively large, as in the period August and September 1964 and 1966, water temperatures were as much as 4°C colder than the corresponding air temperatures. Apparently the larger the release rate through Pepacton's low-level outlet, the longer it takes for climatologic conditions and the natural inflow of tributaries to warm the release water. This same relationship was observed at other sampling sites as far downstream as Callicoon, 59 miles below the reservoir.

It is possible to estimate the effect of releases from Cannonsville Reservoir on water temperatures at Stilesville by assuming that the 1963 relation (fig. 5) and the slope of the regression curves (fig. 4) both approximate the relationship between air and water temperature at this site prior to reservoir storage. By plotting the monthly mean water temperatures at Stilesville against the monthly mean air temperatures at Deposit for the period May-September 1963 and then drawing lines through each point based on the assumed slopes, the relation between these parameters can be estimated (fig. 8). Therefore, natural monthly mean water temperatures can be estimated, based on observed air-temperature data. The difference between the estimated monthly mean water temperatures and the observed temperatures (table 5) is the estimated influence at Stilesville of releases from Cannonsville Reservoir. These estimates are summaried in table 2. Similar estimates were made for the effect of releases from the reservoir on water temperature at Hale Eddy, the effect at Hancock of releases from the Pepacton Reservoir, and the effect at Lordville and Callicoon of releases from both reservoirs (table 2). These estimates were made with

Table 2.—Estimated effect of summer reservoir releases on downstream water temperature (° C) at selected sites in the upper Delaware River basin

Site			1964					1965		
Site	М	J	J	A	s	M	J	J	A	8
Stilesville Hale Eddy Hancock Lordville Callicoon	$-\overline{1}$	-6 	$-5 \\ -2$	$-\frac{0}{6}$	$-1 \\6$		$-2 \\ -1$	$-3 \\ -1$	$ \begin{array}{r} -1 \\ -1 \\ 0 \\ 0 \\ 0 \end{array} $	$ \begin{array}{r} -2 \\ -3 \\ -4 \\ -4 \\ -2 \end{array} $
Site			1966					1967		
	M	J	J	A	8	M	J	J	A	٤
Stilesville Hale Eddy Hancock Lordville Callicoon		-2	-9	$+1 \\ -4$	$-\frac{0}{7}$	-3	-4 	-5	-8 -5	-7 -(-1

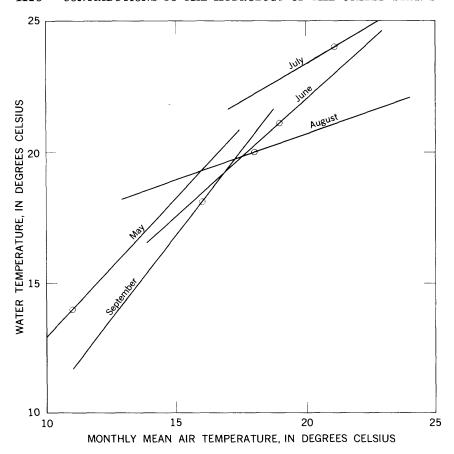


FIGURE 8.—Estimated relation of monthly mean stream and air temperatures at Stilesville, May-September 1964-67, assuming no reservoir effect.

the assumption that natural monthly mean water temperatures for all downstream sampling sites would be 1°C warmer than that of the observed 1963 water temperatures at Stilesville, based upon previously noted observations of monthly mean water temperature between Narrowsburg and Barryville.

Analyses of monthly air and water temperatures, streamflows, and depth of reservoir storage indicate that the influence of the Cannonsville and Pepacton Reservoir releases on downstream water temperature is largely dependent upon five factors:

- 1. Presence of thermal stratification in the reservoir.
- 2. Depth at which water is released from the reservoir, that is, whether by low-level outlet or by spilling over the reservoir's crest.
- 3. Rate of release from the reservoir.

- 4. Distance downstream from the reservoir.
- 5. Concurrent climatologic conditions.

Monthly mean ranges of water temperatures from May through September 1963–67 are reported in table 5. Ranges at Stilesville were as low as 1°C in July 1966 and as high as 7°C for June 1966; those at Hale Eddy were as low as 4°C in August 1966 and as high as 8°C for May 1965. Monthly mean ranges at Hancock on the West Branch Delaware River and at the other sites on the main stem are mostly between 2° and 3°C.

Figure 9 shows that the monthly mean ranges in air temperature at Deposit were more than twice that of monthly mean ranges in water temperature at Stilesville during the period May-September 1963-67. During 1963, the monthly mean minimum water temperatures at Stilesville were always higher than the corresponding mean minimum air temperatures, whereas during 1964-67, the mean minimum water temperatures were at times less than the mean minimum air temperatures.

Figure 9 also shows that large releases of water, presumably from the hypolimnion, decreased both the maximum and minimum observed water temperatures and, thus, shortened the range of temperature. Small releases decreased the monthly minimum, probably had less effect on maximums, and, thus, increased the range. Note that releases of warmer water, as in August and September 1964–66, tended to approximate the natural relation observed between the air and water temperature ranges in 1963. By comparing data for May 1963 with those for 1967, one observes that large spills over the dam in 1967 appeared to have lowered both maximum and minimum water temperatures in relation to corresponding air temperatures and to have greatly contracted the range.

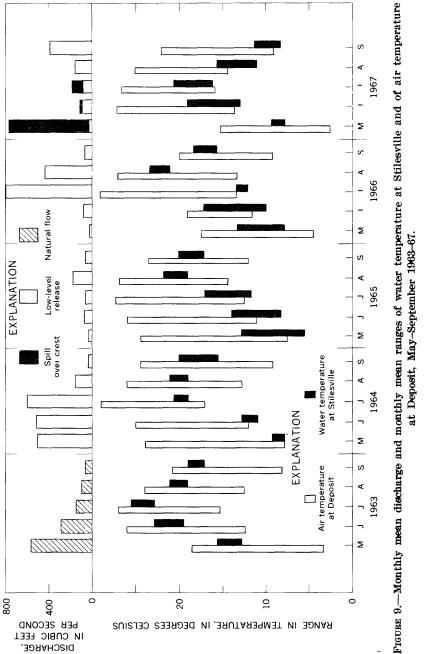
A similar comparison of the monthly mean ranges in water temperature at Hale Eddy with those of air temperature at Deposit (fig. 10) suggests that the rate of release from the reservoir (fig. 9) does not affect ranges at Hale Eddy as much as those at Stilesville. The ranges shortened to 4°C in August and September 1966, apparently as a result of large releases of warmer water.

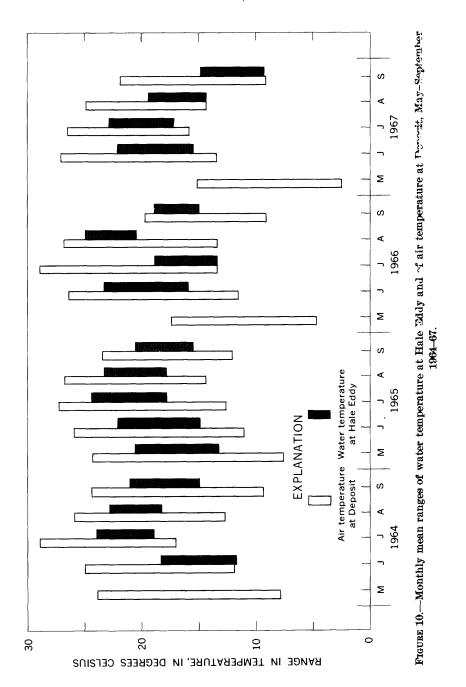
Monthly mean ranges at Hancock on the East Branch Delaware River and at the other sites on the main stem are at most only slightly affected by Pepacton Reservoir releases.

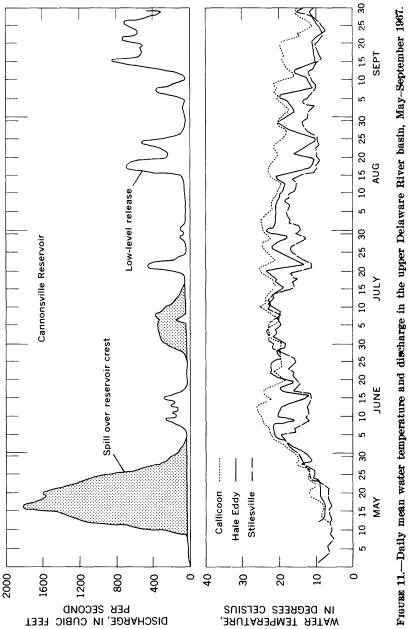
DAILY WATER TEMPERATURES

The varied effects of releases from the low-level outlet and water spilling over the Cannonsville Dam on the water temperatures at Stilesville, Hale Eddy, and Callicoon from May to September 30, 1967, are illustrated in figure 11. During the entire period, water was released

DISCHARGE,







through the low-level outlet; water spilled over the dam from May 10 to June 5 and from June 27 to July 14. No temperature data were obtained at Hale Eddy prior to May 12. From May 12 to September 30, excepting three periods (May 12–23, June 4–5, and September 22–2°), the river water was colder at Stilesville than at Hale Eddy when water did not spill over the dam, and warmer at Stilesville than at Hale Eddy when water did spill over the dam. This suggests that the reservoir was stratified during most of the period. In addition, the effects observed during May 12–23 and September 22–26 suggest the insignificance of thermal stratification in relation to climatologic conditions during May and late September. On June 4 and 5, the volume of water released through the low-level outlet approximated the volume that spilled over the dam. Thus, because of mixing, the water that spilled over the dam did not influence downstream temperature as expected.

The effect of water spilling over the dam on water temperature at Stilesville was quite evident. For example, water temperatures at Stilesville increased 7°C during the period June 26 to July 2 as spill over the dam increased from 0 to 320 cfs and then decreased 7°C as spill over the dam decreased back to zero by July 15. The daily mean water temperatures at Stilesville were consistently higher when there vas spill over the dam and lower when water was released only through the low-level outlet.

Spilling apparently had little or no effect on water temperatures at Hale Eddy and Callicoon. In contrast, temperatures decreased markedly as far downstream as Callicoon, when the flow past Stilesville was augmented by releases through the low-level outlet. For example, when the flow past Stilesville was increased from 29 cfs on August 14 to 679 cfs on August 18 (fig. 11), water temperature decreased 8°C at Stilesville, 10°C at Hale Eddy, and 3°C at Callicoon.

Data collected from May to September 1964–66 show that releases through the low-level outlet from Cannonsville and Pepacton Reservoirs sometimes were accompanied by decreases in water temperatures as far downstream as Narrowsburg. In 1966, releases through the low-level outlet at Cannonsville Reservoir lowered water temperatures by 13°C (Williams, 1968, incorrectly reported 26°F) at Hale Eddy, 4°C at Callicoon, and 1°C at Narrowsburg (Williams, 1968). Releases through the low-level outlet at Pepacton Reservoir lowered water temperatures by 11°C at Hancock, 3°C at Callicoon, and 1°–3°C at Narrowsburg.

To confirm that a change in the proportionality between climatologic conditions and water temperature had occurred at the time reservoir releases began, double-mass curves were constructed. A double-mass curve for June 19-30, 1966, is given in figure 12 as an example. The

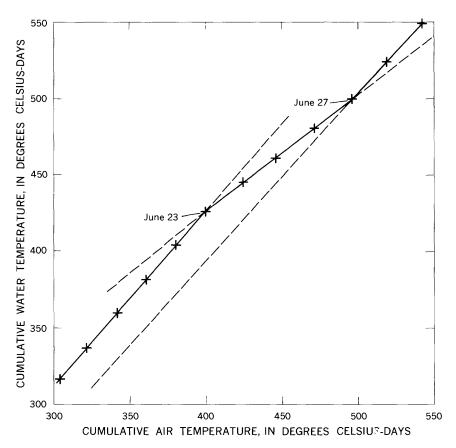


FIGURE 12.—Double-mass curve—cumulative air temperature at Deposit versus cumulative water temperature at Hale Eddy, June 19-30, 1966.

cumulative air-temperature values at Deposit were plotted against the cumulative water-temperature values at Hale Eddy. Streamflow at Stilesville increased from 23 cfs on June 22 to 420 cfs on June 24, 1966, then decreased on June 25. Although air temperatures at Deposit increased 4°C, water temperatures at Hale Eddy, 8.1 miles downstream from the dam, decreased 7°C. When releases were cut back on June 25, water temperatures at Hale Eddy rose 12°C whereas air temperatures decreased about 2°C. The change in slope on June 23 indicates a change in proportionality between climatologic conditions and water temperatures. This suggests that the rate of increase of water temperature was inversely proportional to the rate of increase of air temperature. As water releases were reduced (June 27), the slope approximated that defined prior to the start of releases. Dashed lines, representing ex-

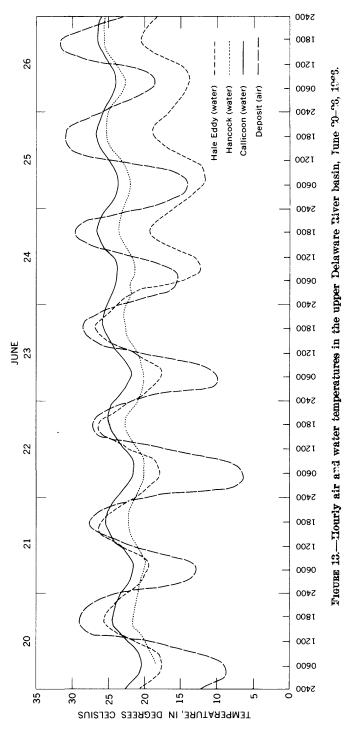
tension of defined linear relations, have been added to this illustration to facilitate its interpretation.

To examine the diurnal fluctuation in temperature, plots were prepared of hourly water temperatures at Hale Eddy, Hancock, and Callicoon and compared with air temperatures at Deposit. The combined plot of water and air temperatures for June 20–26, 1966, is shown in figure 13. The temperatures plot as sinusoidal curves, with the curve for Hale Eddy having the largest amplitude of the three water-temperature curves. The maximum daily water temperature occurred between 4:00 p.m. and 6:00 p.m., within an hour of maximum air temperature. The minimum water temperatures generally occurred between 6:00 a.m. and 9:00 a.m., about 2–3 hours after the minimum air temperature.

Releases from Cannonsville Reservoir were steady at 23 cfs until 3:00 p.m. on June 23; the release rate then increased to 400 cfs by 3:00 p.m. on June 25. The axis of the curve for Hale Eddy dropped about 7°C on June 24 with a slight decrease in the range of diurnal fluctuations. This was in response to increased volumes of cold water at the same time that air temperatures rose slightly. The large diurnal fluctuation is probably related to the close proximity of Hale Eddy to Cannonsville Reservoir. Water released from the reservoir during the morning receives maximum solar radiation while in transit to Hale Eddy, whereas water released during the evening receives little or no solar radiation. Before reaching Callicoon, water from the reservoirs was affected by both day and night conditions, accounting for the lower daily minimum water temperatures. The larger volume of water in the East Branch Delaware River above Hancock (384–578 cfs) reacts less to diurnal temperature variations.

The effects on water temperature of increased releases on Jure 23 were observed at Callicoon on June 25. Water temperatures at Callicoon were usually about 3°C above those at Hancock on the East Branch Delaware River prior to the morning of June 25. Temperatures at Callicoon stabilized on June 25, whereas temperatures at Hancock, which were unaffected by releases from Cannonsville Reservoir, increased to within 1°C of those at Callicoon by late in the evening. Although there was no decrease noted in water temperature at Callicoon, there was a change in its rate of increase. This change indicated that the effects of releases from Cannonsville Reservoir were shadowed by climatological conditions.

The preceding discussion indicates that water temperatures ir the West Branch Delaware River between Cannonsville Reservoir and the confluence with the East Branch are dependent upon the temperature of the water released from the reservoir, the rate of release, the



downstream distance from the reservoir of the site at which water temperature observations are being recorded, and variations in climatologic conditions during the period of observation.

Similar analyses of daily water temperatures in the East Branch produced similar conclusions concerning the influence of releases from the Pepacton Reservoir upon temperature observations at downstream sampling sites. The estimated traveltime for the effect of increased releases from this reservoir on water temperatures at Hancock, based on observations made during August 1964 and 1966 was about 1.5 days. Similarly, the effects of these releases were observed at Callicoon in about 2 days.

Increased release rates from the Cannonsville Reservoir were observed to have a greater effect on daily maximum water temperatures at Stilesville than on daily minimum water temperatures (fig. 14). The release rate during June was about 25 cfs except on June 10, when the rate was 37 cfs and on June 23–30, when the rate ranged between

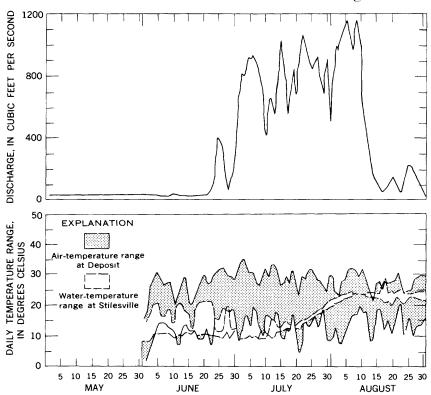


FIGURE 14.—Water-temperature range and discharge at Stilesville and air-temperature range at Deposit, June-August 1966.

51 and 420 cfs. Minimum water temperatures during June ranged between 8° and 11°C, whereas maximum water temperatures ranged between 11° and 21°C. Maximum and minimum water temperatures related well to maximum and minimum air temperatures until June 23, when release rates were increased greatly. Subsequent to this date and throughout July and early August daily maximum water temperatures responded to the variations in the release rate. However, during this period, daily minimum water temperatures showed little or no response to corresponding air temperature variations or to reservoir-release rates. After August 15, when release rates were reduced to about 140 cfs, the response to variations in air temperature were again observed. The higher release rate diminished the effect of climatologic conditions on the water temperature.

Analyses of daily minimum water temperatures observed in 1964 at Stilesville suggest that water released after the end of June was from the thermocline and epilimnion. In 1965, data suggest that releases after mid-July were from the thermocline and, by August, from the epilimnion. Water released through the low-level outlet during 1967 was probably entirely from the hypolimnion.

The influence of releases from Cannonsville Reservoir on daily maximum and minimum water temperatures at Hale Eddy was dependent upon the magnitude of the releases. For example, when releases increased from 23 to 420 cfs during June 23–25, 1966, the daily maximum water temperature decreased 10°C at Stilesville and 8°C at Hale Eddy, and the daily minimum water temperature, which was not affected at Stilesville, dropped 6C at Hale Eddy. When releases increased from 51 to 915 cfs between June 28 and July 4, the effect on both maximum and minimum water temperatures was greater at Hale Eddy than it was at Stilesville, these releases decreased both the maximum and minimum water temperatures 13°C at Hale Eddy. The maximum water temperatures at Stilesville decreased 12°C, whereas the minimum water temperatures decreased only 3°C.

The influence of large releases on the daily minimum water temperatures was more apparent at Hale Eddy than at Stilesville. Previous small releases had lowered the daily minimum at Stilesville near to the temperature of the water released. Therefore, a large release did not lower the daily minimum as much at Stilesville as it did at Hale Eddy.

Further analyses of the data indicated releases from Cannonsville Reservoir decreased the daily maximum water temperatures more than the daily minimum temperatures in the West Branch Delaware River as far downstream as Hale Eddy. The effect upon the daily maximum water temperatures in the mainstem Delaware River was only slightly greater than that on daily minimums. Releases from Pepacton

Reservoir similarly lowered daily maximum water temperatures on the East Branch Delaware River at Hancock by about 1°C more than daily minimums and were only slightly more effective on daily maximum water temperatures than daily minimums as far downstream as Callicoon.

FREQUENCY ANALYSES

Frequency analyses (table 3) were performed on hourly water-temperature data at each site for the months of June, July, and August. A comparison of monthly temperature frequencies at a site can be used to indicate which period is most thermally suitable for various species of aquatic life at that site. Also, a comparison of monthly stream-temperature frequencies at one site with those at other sites can be used to indicate which site is most thermally suitable. For example, if the range of thermal tolerance of a particular species is between 16°C and 29°C, then the frequencies reported in table 3 indicate that July is the most thermally suitable month at Hancock, that Stilesville and Hale Eddy are not suitable, and that the most suitable sites are downstream from Callicoon.

CONCLUSIONS

The objectives of this cooperative investigation were to (1) define the variations in water temperature during the summer months, (2) determine the effect of reservoir releases on downstream water temperatures, and (3) provide data from which the thermal suitability of the study area as a spawning and nursery habitat for anadromous fish can be determined.

The effect of climatologic conditions on stream temperatures in the upper Delaware River basin is to increase the water temperatures from February through July and to decrease temperatures from August through January. Water temperatures of the Delaware River between Narrowsburg and Barryville normally relate well to air temperature at Port Jervis. On the East and West Branches and on the main stem Delaware River, as far downstream as Callicoon, the predominant controlling influence on stream temperature seemed to be reservoir releases. This influence decreased with distant downstream, but was a major factor as far downstream as Callicoon. The influence that releases from Cannonsville and Pepacton Reservoirs have on downstream water temperatures was dependent upon five factors:

- 1. Thermal stratification in the reservoir.
- 2. Depth at which water was released.
- 3. Rate of release.

- 4. Distance downstream from the dams.
- 5. Climatologic conditions.

The effect of relatively large releases from both reservoirs on daily mean water temperature was observed as far downstream as Narrowsburg, 71.0 miles below Pepacton Reservoir and 55.9 miles below Cannonsville Reservoir. Releases from Cannonsville Reservoir were observed to decrease daily mean water temperature by 13°C on the West Branch at Hale Eddy, 8.1 miles downstream, and by about 1°C on the Delaware River at Narrowsburg, 55.9 miles downstream. Releases from Pepacton Reservoir have been observed to decrease daily mean water temperature by 11°C on the East Branch Delaware River at Hancock, 31.0 miles downstream, and between 1°-3°C at Narrowsburg, 71.0 miles downstream.

When Cannonsville Reservoir had less than 50 percent of its total capacity, as during most of the study period, 1964–66, releases through the low-level outlet from May through July were thought to have drained the hypolimnion. Thus, warmer water from the upper strata or epilimnion was released during August–September. Water spilling over the reservoir's crest, beginning in 1967, was observed to have similar effects on downstream water temperatures as low-level releases from the epilimnion.

Releases from Cannonsville Reservoir decreased the daily-maximum water temperatures more than the daily-minimum temperatures in the West Branch Delaware River as far downstream as Hale Eddy. The effect upon the daily maximum water temperatures in the main stem Delaware River was only slightly greater than that on daily minimums. Releases from Pepacton Reservoir similarly lowered daily maximum water temperatures on the East Branch Delaware River at Hancock by about 1°C more than daily minimums and was only slightly more effective on daily maximum water temperatures than daily minimums as far downstream as Callicoon.

Relatively large monthly releases from Cannonsville Reservoir decreased the monthly mean range to as little as 1°C on the West Branch Delaware River at Stilesville. In contrast, relatively small releases (less than 25 cfs) produced a monthly mean range of about 7°C. Monthly mean water temperature ranged between 4° and 8°C ir the West Branch Delaware River at Hale Eddy. In the East Branch Delaware River at Hancock and at other sites on the main stem, monthly mean temperatures ranged mostly between 2° and 3°C. Little or no effect from releases upon the monthly mean range was observed at these latter sampling sites.

Table 3.—Summary of frequency analyses of water temperature, in degrees Celsius, in the upper Delaware River basin, 1964–67

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1-4250.00	1-4250.00 West Branch Delaware River at Stilesville, N.Y.	Aug. 1964–66 June 1964–67 July 1964–67	,					-						2220
1-4265.00	1-4265.00 West Branch Delaware River at Hale Eddy, N.Y.	Aug. 1964–67. June 1964–67. July 1964–67.			`									0 ∞ ۲- 2
1-4272.07	Delaware River at Lordville, N.Y.	Aug. 1964–67 June 1965–67 July 1965–66]∞2i
1-4274.90 Del	Delaware River at Callicoon, N.Y.	Aug. 1965-66 June 1964-67 July 1964, 1966-67								-				254
1-4277.40	Delaware River at Narrowsburg, N.Y.	Aug. 1964–67 June 1965–67 July 1965–67												922
$^{1-4285.00}$ Del	Delaware River above the Lackawaxen River near	Aug. 1965–66 June 1964–57 July 1964–67												17
1-4321.60	Barryville, N.Y. 1-4321.60 Delaware River at Barryville, N.Y.	Aug. 1964–66 June 1964–66 July 1964–67 Aug. 1964–66	27 27 27 27 27 29 28 28 28 28 26 28		222022 22022 2022 2022 2022 2022	255 24 25 25 25 25 25 25 25 25 25 25 25 25 25	22224		55555555555555555555555555555555555555		21 21 19 18 17 16 21 20	10081 10081		0119 184 184
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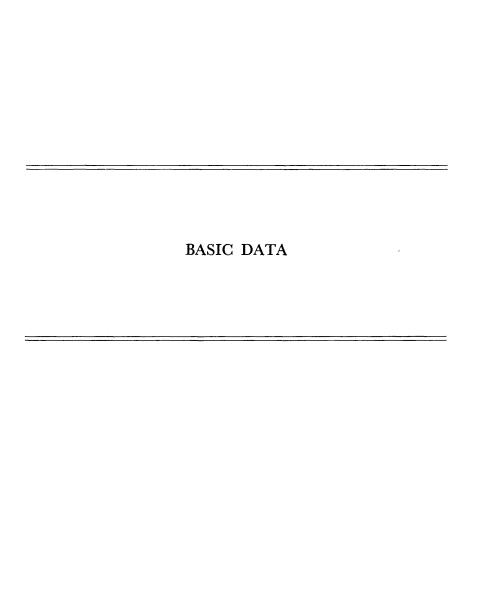


Table 4.—Summary of water temperature records, in degrees Celsius, in the upper Delaware River basin, May-September 1963-67

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Table 4.—Summary of water temperature records, in degrees Celsius, in the upper Delaware River basin, May-September 1963-67.—Con.

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Table 4.→Summary of water temperature records, in degrees Celsius, in the upper Delaware River basin, May-September 1963-67.—Con.

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Table 4.—Summary of water temperatur e records, in degrees Celsius, in the upper Delaware River basin, May-September 1963-67.—Con.

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	18		118 117 118	12021338		17 16 20 20 23 23 23 23 23 21 18	4422	21233332	83	325
	17		118 117 118	12881888		119 119 118 22 22 24 24 18	4222	147233344		325
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	13	Barryville,	24 118 128 138	222222		22 22 23 24 24 25 25 25 25 25 25 25 25 25 25 25 25 25		173427		ກຊ
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	6	River	819618	888888		22 23 23 23 23 23 23 23 23		488888		
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	3	21.6	41 113 113 114	222222		13 20 20 20 20 20 20 20 20 20 20 20 20 20	= ×88	12223823		
	4	01-4321.60	118	222222		118 116 120 120 130 130 130	116 21	288838		
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	23		20	222222		116 113 118 118 221 221 129 119		888888	- 1	
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	Day		May, max min June, max	July, max. August, max. September, max. min		May, max June, Max June, Max July, max min August, max September, max min	May, maxJune, max	July, max. August, max. September, max. min.	May, max	June, max
-	Ď	l	M£ Ju	Ju At		Ma Ju Ju Au Se,	Mi	Ju Au Sej	Ψ	Ju

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July, max
min
August, max
min
September, max

Table 5.—Summary of monthly statistics of water-temperature data, in degrees Celsius, collected in the upper Delaware River basin, May-September 1963-67

Date		ean mum	Mean minimum	Mini- mum	Monthly mean	Maximum range	Mean range	Minimum range
	1–4215.	East	Branch Dela	ware Rive	er at Hanco	rk, N.Y.		
1964								
May	24	19. 4	17. 2	14	18	5	2, 2	
June	29	26. 1	23. 3	16	24	4		
July August September	19	16. 7	25. 5 15. 0	13	15	3	2. 8 1. 7	
1965	00	10.0	16. 7	10	18	6	2. 2	
May June	23 27	18. 9 21. 7	18. 3	11	20	5	3. 4	
July	27	23. 9	21. 1	18	23	5	2. 8	
August	28	23. 3	21. 1	15	22	4	2, 2	
September	25	19. 4	17. 2	12	18	5	2. 2	
1966		10.0		•	10			
May	18 27	12. 8 21. 1	11. 1 18. 9	6 13	$^{12}_{20}$	4 6	1.7	
June July	27 28	24. 4	18. 9 22. 8	21	20 24	0 4	2. 2 1. 6	
August	26 26	18. 3	16. 7	12	18	5	1.6	
September	17	11, 1	10. 0	8	10	3	1. 1	
1967								
May								
June								
July								
September	22	18. 3	16. 1	11	17	5	2. 2	•
	1-4250.	Wes	t Branch Dela	ware Riv	er at Stilesv	ille, N.Y.		
1963 May	19	15, 6	12. 8	9	14	6	2. 8	
June	28	22. 8	19. 4	15	21	5	3. 4	
July	31	25. 6		18	24	6	2.8	
August	25	21. 1	18. 9	15	20	š	2. 2	
September	23	18. 9	17. 2	13	18	5	1. 7	
1964								
Мау	14	9, 4	7.8	6	9	4	1.6	
June	17	12.8	11.1	8	12 20	5 4	1.7	
July	24 24	20. 6 21. 1	18. 9 18. 9	16 18	20 20	4	$\begin{array}{c} 1.7 \\ 2.2 \end{array}$	
August September	24 24	20. 0	15. 6	12	18	8	4.4	
1965								
Мау	18	12.8	5. 6	1	9	11	7. 2	
June	19	13, 9	8.3	7	11	11	5.6	
July	24	17. 2	11.7	.8	14	11	5. 5	
August	26	21.7	18.9	14	20	6	2. 8 2. 8	
September	25	20.0	17. 2	13	19	6	2. 8	
1966 May	19	13. 3	7.8	4	11	11	5 5	
June	21	17, 2		8	11	11	5. 5 7. 2	
July	21	13. 3		9	13	6	1. 1	
August	28	23. 3		19	22	8	2. 2	
	25	18. 3	15. 6	11	17	6	2. 7	
September					9	9		
September								
September 1967 May	17	9.4		4			1.6	
September	24	18.9	13	10	16	12	5. 9	
September 1967 May	17 24 24 21		13 16					

Table 5.—Summary of monthly statistics of water-temperature data, in degrees Celsius, collected in the upper Delaware River basin, May-September 1963-67—Con.

Date	Maxi- mum	Mean maximum	Mean minimum	Mini- mum	Monthly mean	Maximum range	Mean range	Minimum range
	1-4	265. West	Branch Dela	ware Rive	er at Hale E	ddy, N.Y.		
1964					·			
May								
June July	25 29	18. 3 23. 9	11. 7 18. 9	7 16	15 21	9 7	6. 6 5. 0	;
August	26	22. 8	18.3	15	21	8	4.5	
September	27	21, 1	15. 0	10	18	9	6. 1	
1965								
Мау	26	20, 6	13. 3	7	17	10	7. 3	
June	27	22. 2	15.0	11	19	ii	7. 2	
fuly	30	24. 4	17.8	14	21	11	6.6	
August	27	23. 3	17.8	11	21	8	5. 5	
September	27	20. 6	15. 6	9	18	8	5 . 0	
1966								
May								
une	28	23. 3	16. 1	12	19	10	7. 2	
July	25	18.9	13. 3	10	20	10	5.6	;
August	27 26	25. 0 18. 9	20. 6 15. 0	19 9	17 12	8 6	4. 4 3. 9	
September	26	18. 9	15. 0	9	12	ь	3. 9	
1967								
May	27			9				
une uly	27 27	22. 2 22. 8	15. 6 17. 2	10	19 20	11 14	6. 6 5 . 6	i
August	2 5	19. 4	14. 4	8	17	8	5.0	
September	$\frac{24}{24}$	15. 0	9. 4	7	12	9	5. 6	2
		1-4272.0	7 Delaware	River at 1	Lordville, N	.Y.		
1964								
Мау								
May June								
une uly								
une uly August				11	15	3	2. 2	
une uly August September				11	15	3	2. 2	
fune	19	16. 1	13. 9					
une uly August September 1965 May	19	16. 1	13. 9 15. 0	9	16	6	2, 8	
rune Fuly Saugust September 1965 May Inne	19 23 26	16. 1 17. 8 21. 1	13. 9 15. 0 18. 3	9 12	16 20	6 5	2. 8 2. 8	
une July August September 1965 May une une	19 23 26 27	16. 1 17. 8 21. 1 23. 3	13. 9 15. 0 18. 3 21. 1	9 12 19	16 20 22	6 5 4	2, 8 2, 8 2, 2	ĺ
une july August September 1965 May une july August	19 23 26	16. 1 17. 8 21. 1	13. 9 15. 0 18. 3	9 12	16 20	6 5	2. 8 2. 8	(
fune fuly August September 1965 May	19 23 26 27 28	16. 1 17. 8 21. 1 23. 3 22. 8	13. 9 15. 0 18. 3 21. 1 20. 6	9 12 19 15	16 20 22 22	6 5 4 5	2, 8 2, 8 2, 2 2, 2	
une	19 23 26 27 28 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2	9 12 19 15 12	16 20 22 22 22 18	6 5 4 5 4	2. 8 2. 8 2. 2 2. 2 1. 7	,
une	19 23 26 27 28 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2	9 12 19 15 12	16 20 22 22 22 18	6 5 4 5 4	2. 8 2. 8 2. 2 2. 2 1. 7	
une	19 23 26 27 28 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2	9 12 19 15 12	16 20 22 22 22 18	6 5 4 5 4	2. 8 2. 8 2. 2 2. 2 1. 7	
une	19 23 26 27 28 24 19 28 26 26 27 28 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2	9 12 19 15 12	16 20 22 22 22 18	6 5 4 5 4	2. 8 2. 8 2. 2 2. 2 1. 7	
une	19 23 26 27 28 24 19 28 26 26 27 28 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9 13. 3 21. 7 20. 6	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2	9 12 19 15 12 7 14 17	16 20 22 22 22 18	6 5 4 5 4	2.8 2.8 2.2 2.2 1.7	
une	19 23 26 27 28 24 19 28 26 26 27 28 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9 13. 3 21. 7 20. 6	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2	9 12 19 15 12 7 14 17	16 20 22 22 22 18	6 5 4 5 4	2.8 2.8 2.2 2.2 1.7	
une	19 23 26 27 28 24 24 19 28 26 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9 13. 3 21. 7 20. 6 20. 6	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2 11. 7 19. 4 18. 9 18. 9	9 12 19 15 12 7 14 17 16	16 20 22 22 22 18 12 21 20 20	6 5 4 5 4 5 4	2.8 2.8 2.2 2.2 1.7	(
une	19 23 26 27 28 24 19 28 26 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9 13. 3 21. 7 20. 6 20. 6	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2 11. 7 19. 4 18. 9 18. 9	9 12 19 15 12 7 14 17 16	16 20 22 22 22 18 12 21 20 20	6 5 4 5 4	2.8 2.8 2.2 2.2 1.7	
une	19 23 26 27 28 24 19 28 26 24	16. 1 17. 8 21. 1 23. 3 22. 8 18. 9 13. 3 21. 7 20. 6 20. 6	13. 9 15. 0 18. 3 21. 1 20. 6 17. 2 11. 7 19. 4 18. 9 18. 9	9 12 19 15 12 7 14 17 16	16 20 22 22 22 18 12 21 20 20	6 5 4 5 4 5 4 2	2.8 2.8 2.2 2.2 1.7 1.6 1.3 1.7 1.7	

Table 5.—Summary of monthly statistics of water-temperature data in degrees Celsius, collected in the upper Delaware River basin, May-September 1963-67

Date	Maxi- mum	Mean maximum r	Mean minimum	Mini- mum	Monthly mean	Maximum range	Mean range	Minimum range
		1-4274.9	Delaware	River at	Callicoon, I	1.Y.	, ,	
1964								
May	24	19. 4	17. 2	13	18	4	2.2	1
June July	26 30	22. 2 25. 6	19. 4 2 3 . 9	16 21	21 25	4 3	2. 8 1. 7 1. 7	1 0
August	24	21. 7	20, 0	16	21	ь	1.7	0
September	23	18. 9	17.8	13	18	3	1.1	0
1 9 65								
May June	28	23. 3	19, 4	15	21	6	3. 9	1
July			19.4	10		0	0. 9	
August Septemper	29 26	23. 9 20. 6	21. 7 18. 3	16 13	23 19	4 5	2. 2 2. 3	1 0
	20	20.0	10.0	10	15	v	2. 0	v
1966								_
May June	21 29	13. 9 23. 3	12. 8 20. 6	8 15	13	3 5	1.1	0
JulyAugust	29	24.4	21, 7	19	22 23	6	2. 7 2. 7 1. 7	i
August September	26 22	22. 8 15. 6	21. 1 13 . 9	18 11	22 15	4 3	1.7 1.7	1 1 1 0
1967	22	10, 0	10. 3	11	10	0	1. 1	v
May June	28	22.8	20. 6	14	22	4	2. 2	0
July August	27	23. 9	21.7	18	22 23	4	2. 2	1
August September	26 23	21. 7 18. 3	20. 0 15. 6	17 10	21 17	3 8	2. 2 2. 2 1. 7 2. 7	0 1
		1-4277.4 De	laware Riv	er at Narr	owsburg, N	T.Y.		
1965								
May	25 28	20. 0	16. 7	11	18	5	3.3	1
June July	28	23. 3 25. 0	19. 4 22. 2	15 20	21 2 4	6 5	3. 9 2. 8 2. 7	1 1 1
August	29	24. 4	22. 2 21. 7	16	23	4	2.7	1
September	26	20, 6	18. 3	13	19	5	2, 3	0
1966								
Мау	20	13. 3	12. 2	7	13	3	1.1	0
June	29 31	23, 3 25, 0	20.6	15	22 24	4 6	2.7	1
July August	26	23. 0 22. 8	22. 8 21. 7	21 19	24 22	3	2. 2 1. 1	0 1 0 0
September	24	16. 1	14.4	9	15	4	1.7	0
1967								
May								
June July	28 27	23. 9 24. 4	21.7 22.2	16 19	23 23	6 4	2. 2 2. 2	0
August								
September								
1-42	85. Dela	ware River a	bove the La	ckawaxe	n River near	Barryville,	N.Y.	
1 9 64								
Мау			19. 4	16	21 2 5	5 5	3. 4 2. 8	1 1
May	27	22. 8 26. 1	23 3					
May June July August	27 29 26	26. 1 22. 8	23. 3 20. 0	19 18	21	4	2.8	1
May June July August	27 29	26. 1	23. 3	19 18 14	21 19	4	2. 8 2. 2	1 1
May	27 29 26	26. 1 22. 8	23. 3 20. 0	18	21	4	2. 8 2. 2	1
May	27 29 26 24	26. 1 22. 8 20. 0	23. 3 20. 0 17. 8	18 14	21 19	4 4	2. 8 2. 2	1
May	27 29 26 24 24 24	26. 1 22. 8 20. 0	23. 3 20. 0 17. 8	18 14 13 16	21 19 19 21	4 4 5 3	2. 8 2. 2	1 1 1 1 0
May June July August September	27 29 26 24	26. 1 22. 8 20. 0	23. 3 20. 0 17. 8	18 14	21 19	4 4	2. 8 2. 2	1

Table 5.—Summary of monthly statistics of water-temperature data, in degrees Celsius, collected in the upper Delaware River basin, May-September 1963-67—Con.

Date	Maxi- mum	Mean maximum	Mean minimum	Mini- mum	Monthly mean	Maximum range	Mean range	Minimum range
1-4285.	Delawar	e River abov	ve the Lakav	vaxen Riv	er near Bar	ryville, N.Y.	Continu	ıed
1966								
May								
Tune	28	22, 8	20.6	16	22	6	2. 2	
July	29	26, 1	23. 3	20	25	4	2, 7	
August	26	24. 4	21. 7	19	23	4	2.7	
September	26	18. 9	16. 7	12	18	4	2, 2	
1967								
May		04.4	01.7			6	2. 7	
June July	28 27	24. 4 24. 4	21. 7 21. 7	16 19	23 23	7	2. 7	
	27 26		20.6	18	23 22	5	2, 7	
August		22.8		13	22	4	2. 2	
September	25	21. 1	18. 3	13	20	4	2. 8	
		1-4321.6	Delaware R	iver at Ba	ryville, N.	7.		
1964								
May	22	18.9	16. 7	11	18	6	2, 2	
une	26	22, 2	20. 0	16	21	6	2. 2	
uly	30	25. 0	22. 8	21	24	Š	2. 2	
August	26	22. 8	21.7	19	22	2	1, 1	
September	25	20. 0	19.4	16	$\frac{22}{20}$	3	0.6	
1965								
May	23	18, 9	16.1	10	18	6	2.8	
une	27	22. 8	20. 0	16	21	5	2.8	
	27	24. 4	21. 1	18	23	5	3. 3	
uly	27	23.3	21.1	17	22	4	2. 3	
August				14	19	4	1.7	
September	25	20.0	18. 3	14	19	4	1. 7	
1966								
May	20	13.9	12.8	8	13	2	1.1	
une	28	22.8	19.4	14	21	$\begin{array}{c}2\\7\\7\end{array}$	3.4	
uly	30	25.6	22.8	21	24	7	2.8	
August	26	23. 9	22, 2	21	23	2	1.7	
September	23	16.7	15. 6	12	16	3	1.1	
1967								
May				· - -				
une				19		4	1.1	
uly	28	23. 3	22. 2		23	4	1.1	
Lugust								